

## Gravity, Geoid and Height Systems 2016

Stochastic modeling of altimetric sea surface height measurements  
— refined AR models from iterative residual analysis

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## 1 Motivation and Introduction

## 2 Parameterization and Methods

## 3 Used Data

## 4 Results of Test Study

## 5 Summary and Conclusions

## Estimation of the DOT from along track altimetry

## Gravity field model



- ★ GOCE-based
- ★ band-limited
- ⇒ full VCM

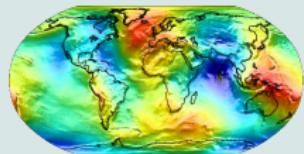
## Altimetry



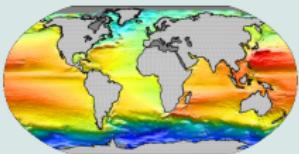
- ★ along-track SSH
- ★ e.g. AVISO Cor-SSH product

separation

⇒ joint adjustment



**Geoid**  
global spherical  
harmonics

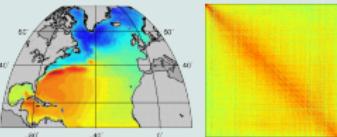


**Dyn. Topography**  
local finite  
elements

**spherical harmonics:**

- ★ here: buffer for high frequency altimetry
- ★ eliminated from NEQ
- ★ future: target quantity

## DT on predefined grid + VCM

**finite elements:**

- ★ target quantity
- ★ currently: local focus
- ★ resolution defined by grid & base functions

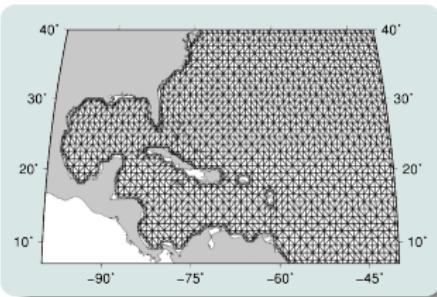
## Gravity Field

- ▶ spherical harmonic base functions, physically motivated
- ▶ global support
- ▶ here up to degree 540

## Dynamic Topography

- ▶ finite element base functions
- ▶ local support
- ▶ here: linear elements, triangulated  $1^\circ \times 1^\circ$  grid (IFEOM), static

here: linear finite elements are set up for a local area (North Atlantic), constant in time  $\rightarrow$  MDT as linear combination of base functions

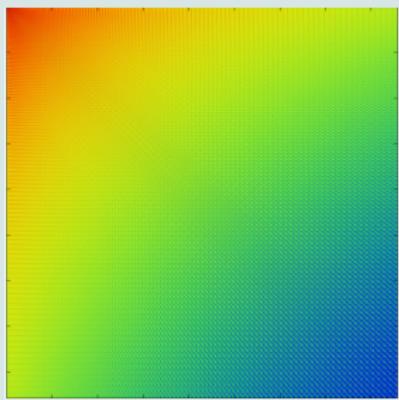


$$\text{MDT}(\theta, \lambda) = \sum_{k \in K} \color{red}{a_k} b_k(\theta, \lambda), \quad \mathbf{x}_{FE} = [\color{red}{a_k}]$$

- ▶ nodal points predefined by ocean model
- ▶ unknowns  $a_k \Rightarrow$  MDT at the nodal points
- ▶ direct access to target quantities

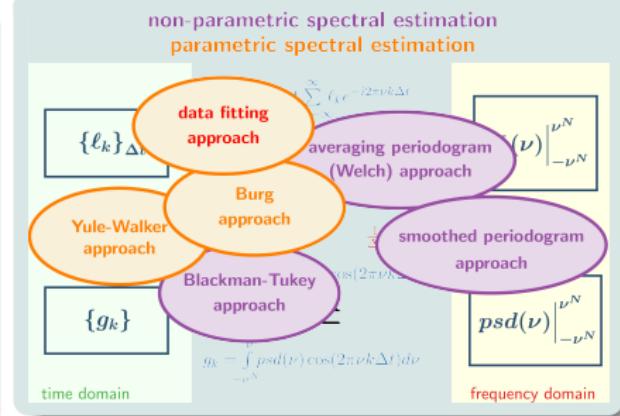
## Gravity Field

- ▶ use of unconstrained models and full covariance matrix:  $\mathbf{x}_{\text{gf}}$ ,  $\Sigma_{\mathbf{x}_{\text{gf}} \mathbf{x}_{\text{gf}}}$
- ▶ NEQs as sufficient statistics of original observations (inversion)
- ⇒ direct use of VCM, no special action required



## Altimetric Observations

- ▶ measurements: time series along the orbit
- ▶ model describes: stochastic measurement errors, model errors
- ⇒ AR( $p$ ) processes iteratively estimated from residuals



- ▶ Robust estimation of AR( $q$ ) process coefficients from residuals  $\mathcal{V}_{\text{ssh}_{m,c,p}}$
- ▶ Iterative refinement: after full adjustment (gravity field and MDT)

$$\mathcal{V}_{\text{ssh}_{m,c,p}}(k) = \mathcal{S}_{m,k} + \mathcal{N}_{m,k}$$

$$\mathcal{S}_k = \sum_{j=1}^q \alpha_j \mathcal{S}_{k-j} + \mathcal{E}_k$$

$\mathcal{S}_k$  ... stochastic process in AR representation

$\mathcal{E}_k$  ... innovation process

$\mathcal{N}_k$  ... additive process (noise)

cf. Schuh et al. (2015)

- ▶  $k\sigma$  rejection in iterative Least squares adjustment
- ▶ AR process coefficients are computed A) *per mission*, B) *per cycle*, C) *per pass*
- ▶ choice of process order  $q$ :  $q = 16$  (A),  $q = 16$  (B),  $q = \text{pass length}/50$  (C)
- ▶ AR process is converted to error covariance function (ecf) via Yule-Walker equation

## 1. Input data and information

- ▶ predefined finite element setup (triangulation/base function type)
- ▶ predefined spherical harmonic setup (maximal resolution)
- ▶ NEQs of gravity field model:  $\mathbf{N}_{\text{gf}}$ ,  $\mathbf{n}_{\text{gf}}$ ,
- ▶ OEQs of ssh observations of a mission  $m$ :  $\mathbf{A}_{\text{ssh}_m}$ ,  $\boldsymbol{\ell}_{\text{ssh}_m}$ , and initial covariance model  $\mathbf{Q}_{\text{ssh}_m}^{(0)}$
- ▶ initial weights for all observation groups ( $w_m$  and  $w_{\text{gf}}$ )

## 2. Compute NEQs for ssh observations

$$\mathbf{N}_{\text{ssh}_m} = \mathbf{A}_{\text{ssh}_m}^T \mathbf{Q}_{\text{ssh}_m}^{(0)-1} \mathbf{A}_{\text{ssh}_m}, \quad \mathbf{n}_{\text{ssh}_m} = \mathbf{A}_{\text{ssh}_m}^T \mathbf{Q}_{\text{ssh}_m}^{(0)-1} \boldsymbol{\ell}_{\text{ssh}_m}, \quad (1)$$

## 3. Combine NEQs (different parameter spaces!)

$$\mathbf{N} = w_{\text{gf}} \mathbf{N}_{\text{gf}} + \sum_m w_m \mathbf{N}_{\text{ssh}_m}, \quad \mathbf{n} = w_{\text{gf}} \mathbf{n}_{\text{gf}} + \sum_m w_m \mathbf{n}_{\text{ssh}_m} \quad (2)$$

4. Rigorous parameter estimation (huge dimensional)  $\tilde{\mathbf{x}} = \mathbf{N}^{-1} \mathbf{n}$
5. Estimate weights for all ( $w_m$  and  $w_{\text{gf}}$ ), continue with 3.
6. Compute SSH residuals  $\mathbf{v}_{\text{ssh}_m} = \mathbf{A}_{\text{ssh}_m} \tilde{\mathbf{x}} - \boldsymbol{\ell}_{\text{ssh}_m}$

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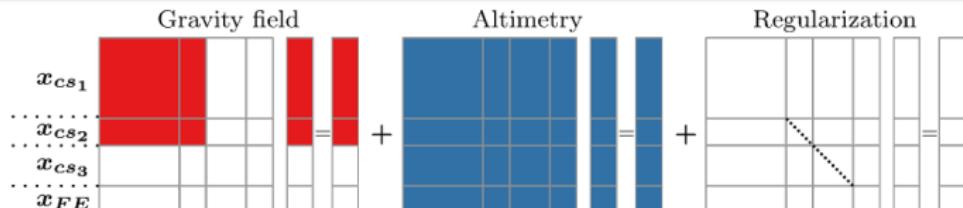
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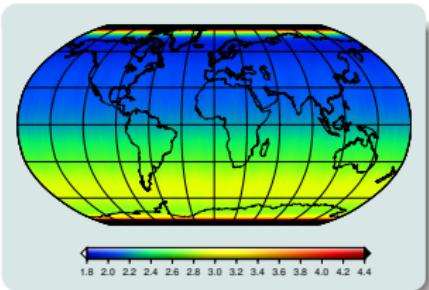
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\* Geoid data: Mainly EGM\_TIM\_RL05 NEQs (Brockmann et al., 2014)

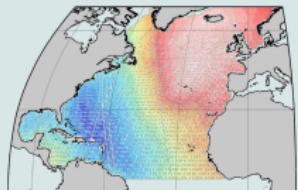
$i$	$l_{\min}$	$l_{\max}$
SGG	2	281
SST	2	150
KAULA	181	540
KAULA	zonal coeff.	



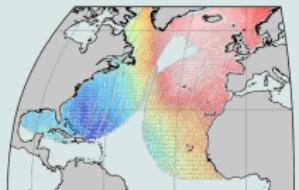
Propagated Geoid  
accuracy of the  
EGM\_TIM\_RL05 model  
at 100 km resolution

\* SSH data:  $\approx$  1 year, 4 missions, mainly from 2007 (AVISO CorSSH)

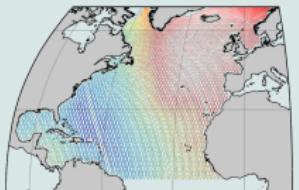
Envisat, cy 53–64,  
1 972 985 obs



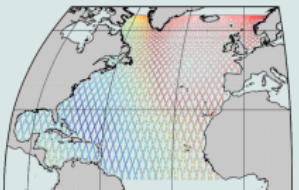
ERS2, cy 121,  
157 512 obs



GFO, cy 184–208,  
1 724 166 obs



Jason1, cy 179–222,  
2 747 694 obs



A) mean analytical ecf  
per mission  $m$   
 $\Rightarrow 4$  analytical ecfs

B) discrete ecf  
per  $m$  & cycle  $c$   
 $\Rightarrow 82$  discrete ecfs

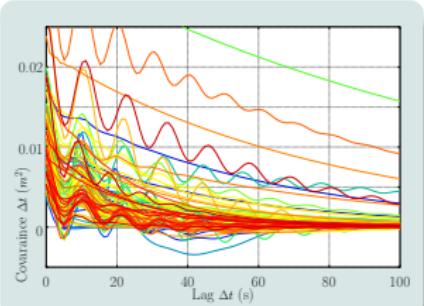
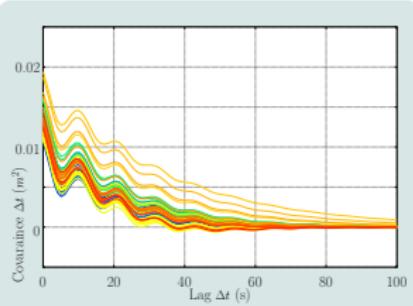
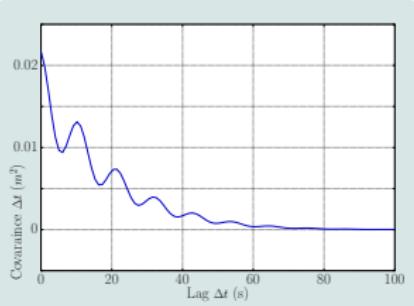
C) discrete ecf  
per  $m, c$  & pass  $p$   
 $\Rightarrow 10k$  discrete ecfs

### Example: Jason 1, error covariance functions

J1,

cycle 179–222

pass 1–84 of cycle 182

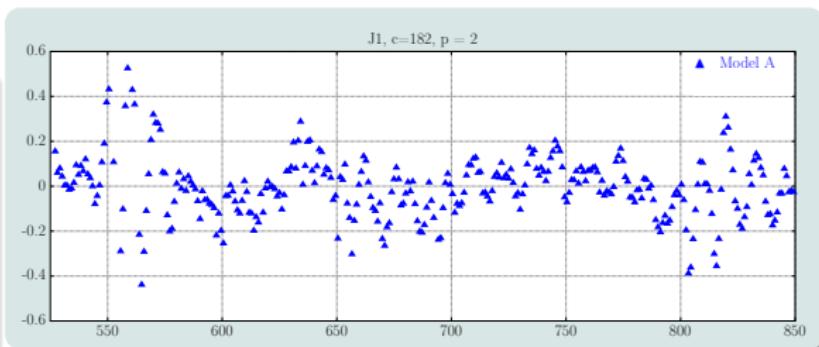
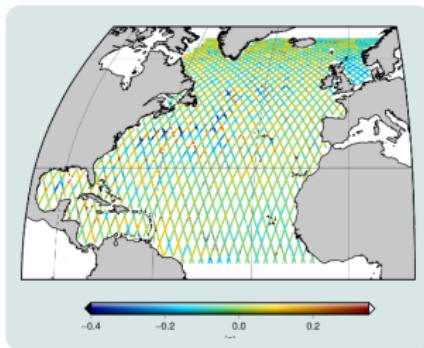


- ▶ mean analytical function

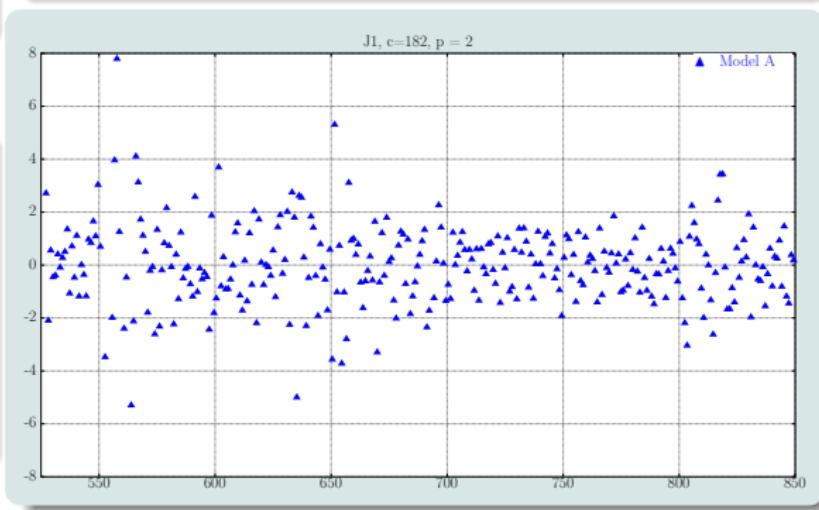
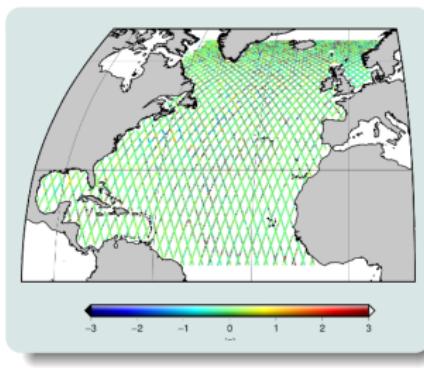
- ▶ stable estimates per cycle
- ▶ changes during cycle?

- ▶ strong variations
- ▶ unstable?
- ▶ AR order changing

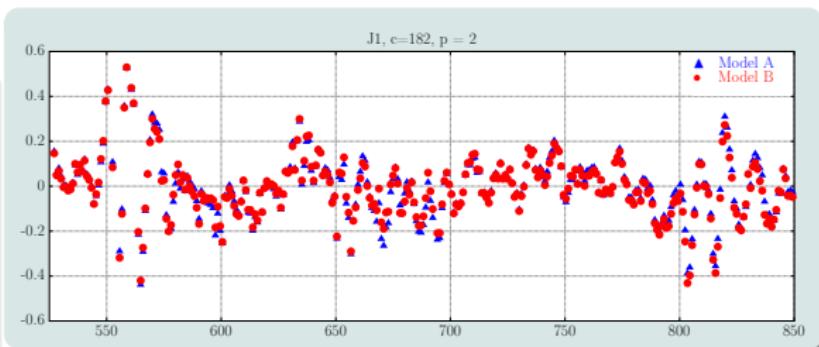
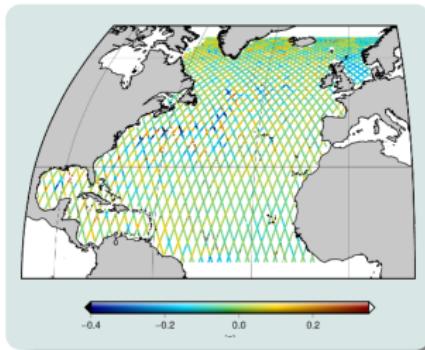
Residuals,  $\sigma = 0.101 \text{ m}$



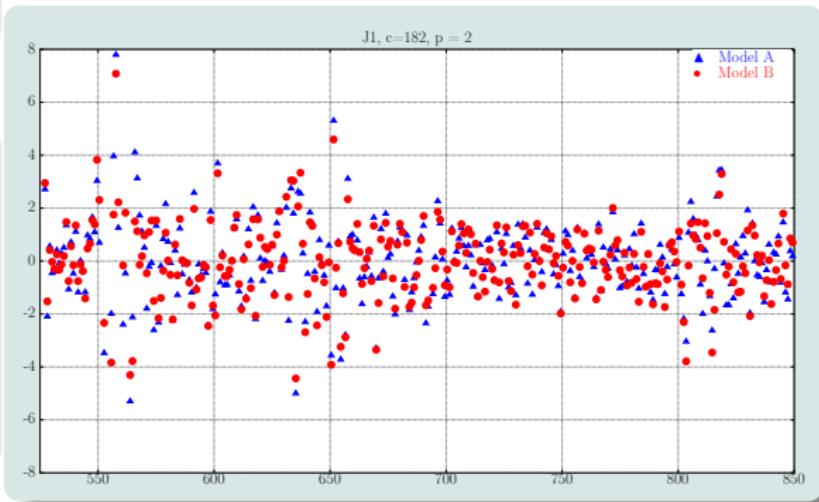
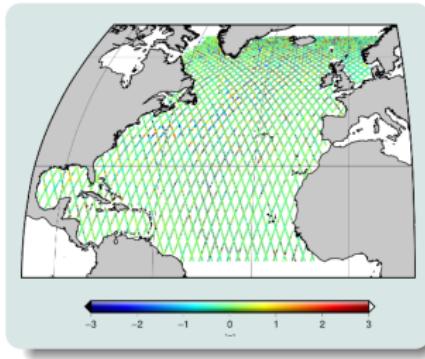
Decorrelated residuals,  $\sigma = 1.16$



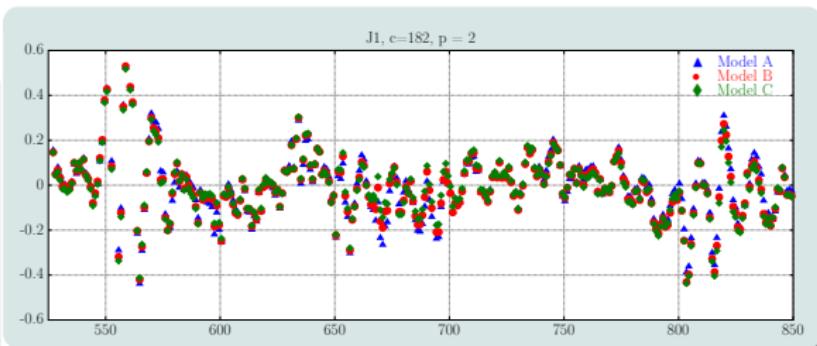
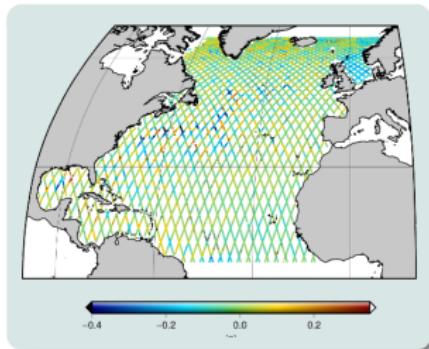
Residuals,  $\sigma = 0.099$  m



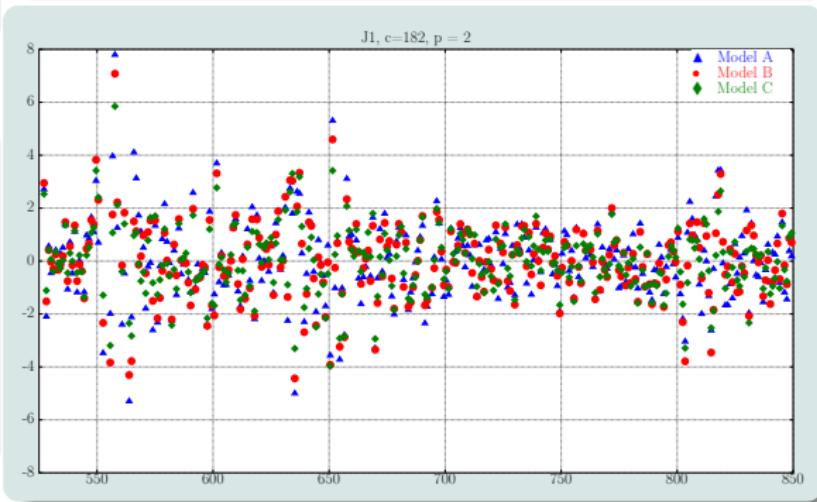
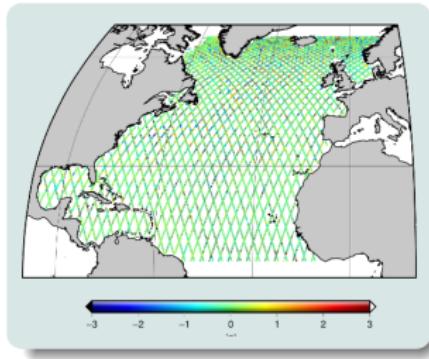
Decorrelated residuals,  $\sigma = 1.10$



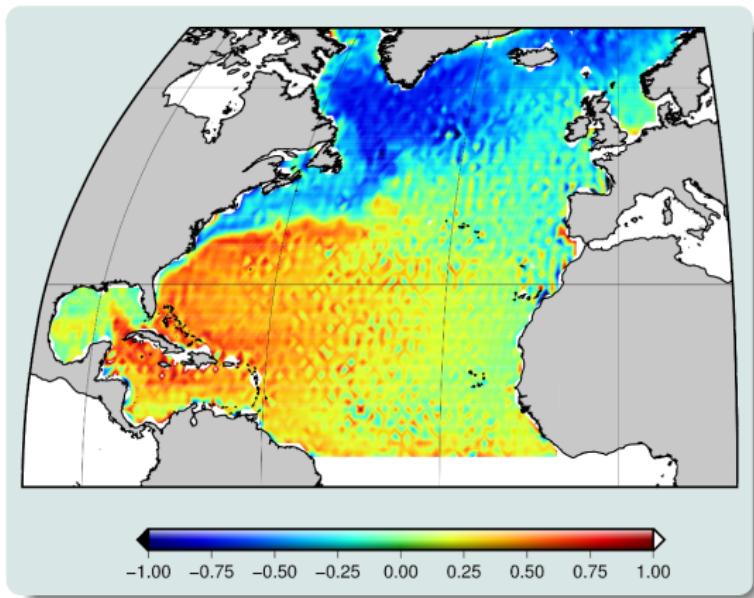
Residuals,  $\sigma = 0.098$  m



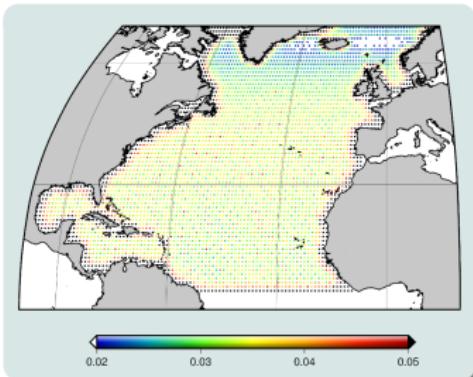
Decorrelated residuals,  $\sigma = 0.96$



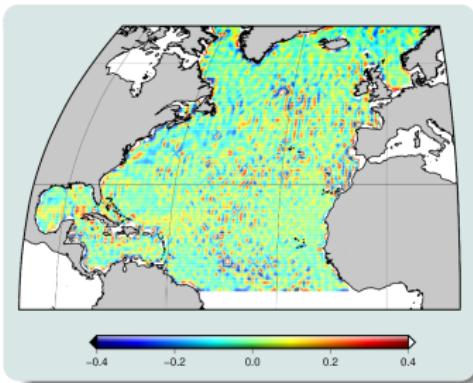
a) MDT model estimate (m)



b) Standard deviation (m)



c) Diff. to CNES-CLS13 (m)

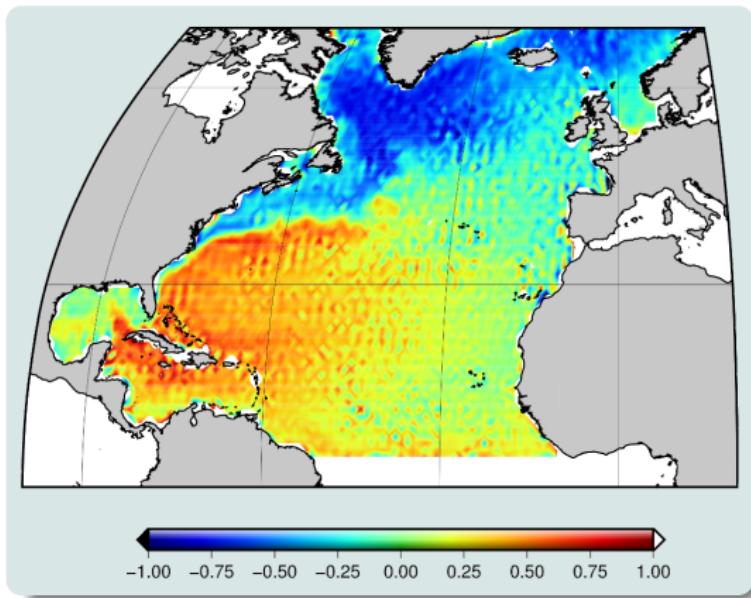


### Statistics (m)

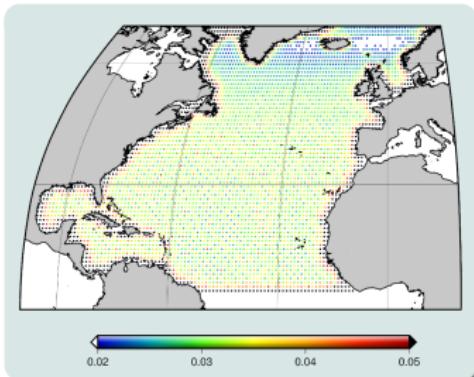
- a)  $\sigma_{\text{rms}} = 0.280 \text{ m}$  (variability)
- b)  $\sigma_{\text{mean}} = 0.053 \text{ m}$  (mean error)
- c)  $\sigma_{\text{rms}} = 0.155 \text{ m}$ ,  $\sigma_{\text{mad}} = 0.148 \text{ m}$

MDT-CNES-CLS13 was produced by CLS Space Oceanography Division and distributed by Aviso, with support from Cnes (<http://www.aviso.altimetry.fr/>)

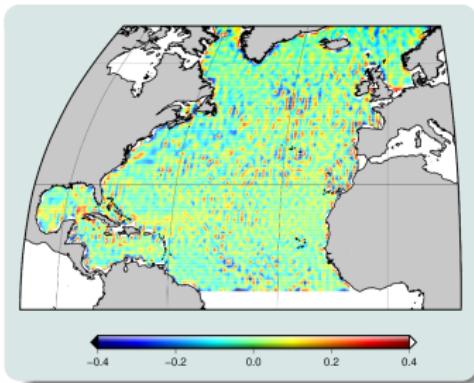
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c) Diff. to CNES-CLS13 (m)

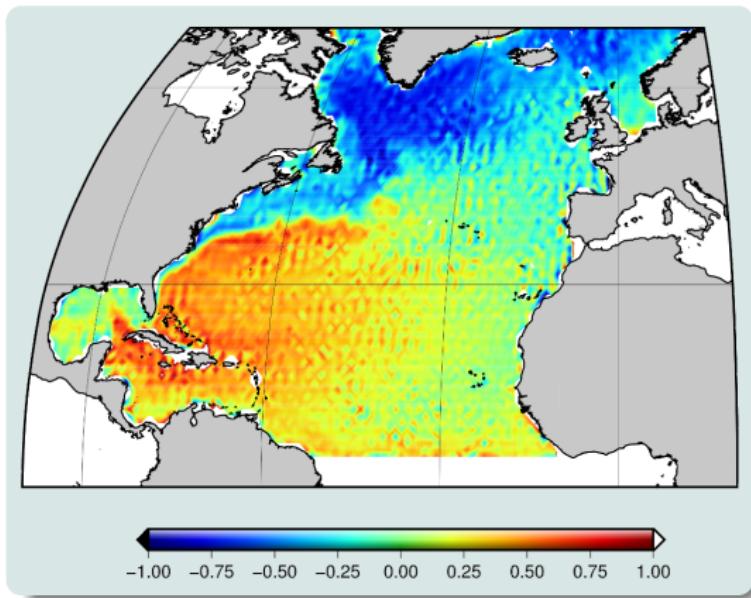


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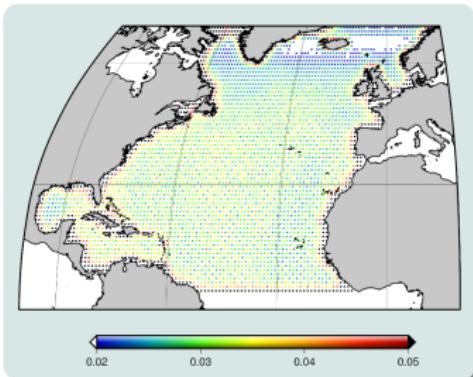
- a)  $\sigma_{\text{rms}} = 0.277 \text{ m}$  (variability)
- b)  $\sigma_{\text{mean}} = 0.051 \text{ m}$  (mean error)
- c)  $\sigma_{\text{rms}} = 0.141 \text{ m}$ ,  $\sigma_{\text{mad}} = 0.135 \text{ m}$

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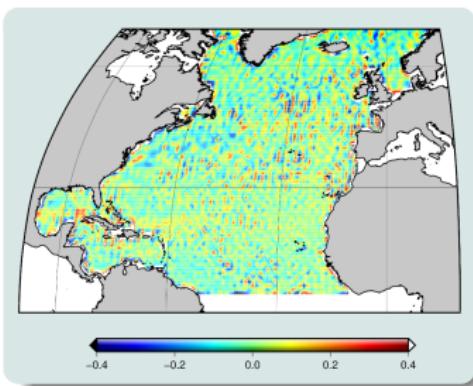
a) MDT model estimate (m)



b) Standard deviation (m)



c) Diff. to CNES-CLS13 (m)



## Summary and Conclusions

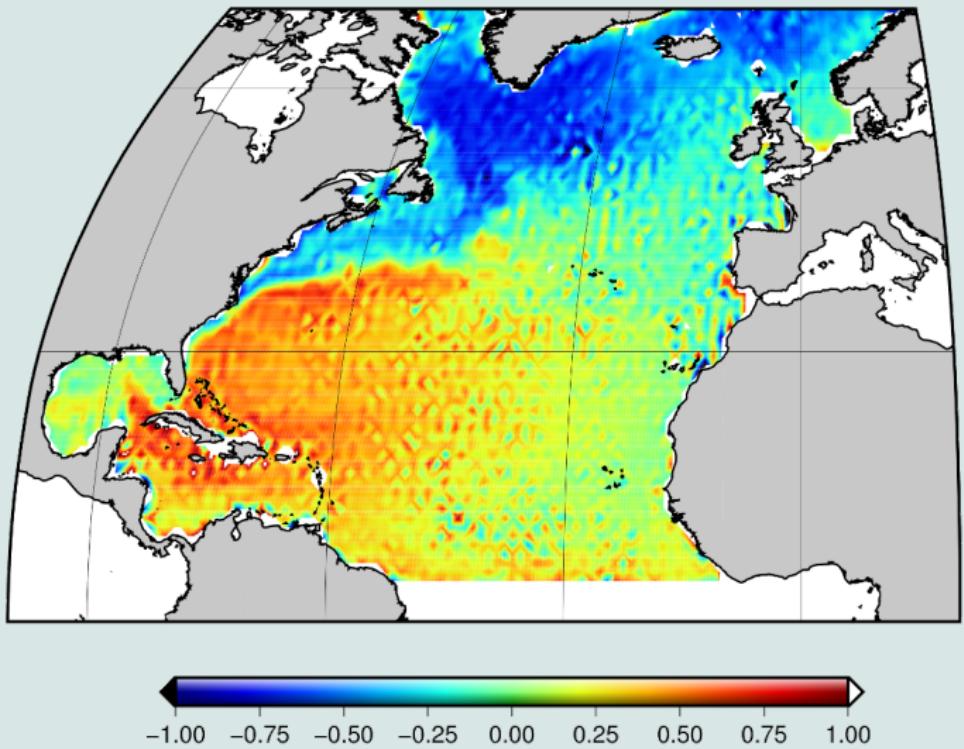
- ▶ AR processes can be used as stochastic model for temporal correlations of along-track SSH observations
- ▶ Conversion to error covariance functions is possible (computational advantages)
- ▶ Estimation can be performed per mission, per cycle and per pass
- ▶ To deal with non-stationarity of residuals: estimation per pass was tested  
⇒ no obvious gain yet

## Outlook

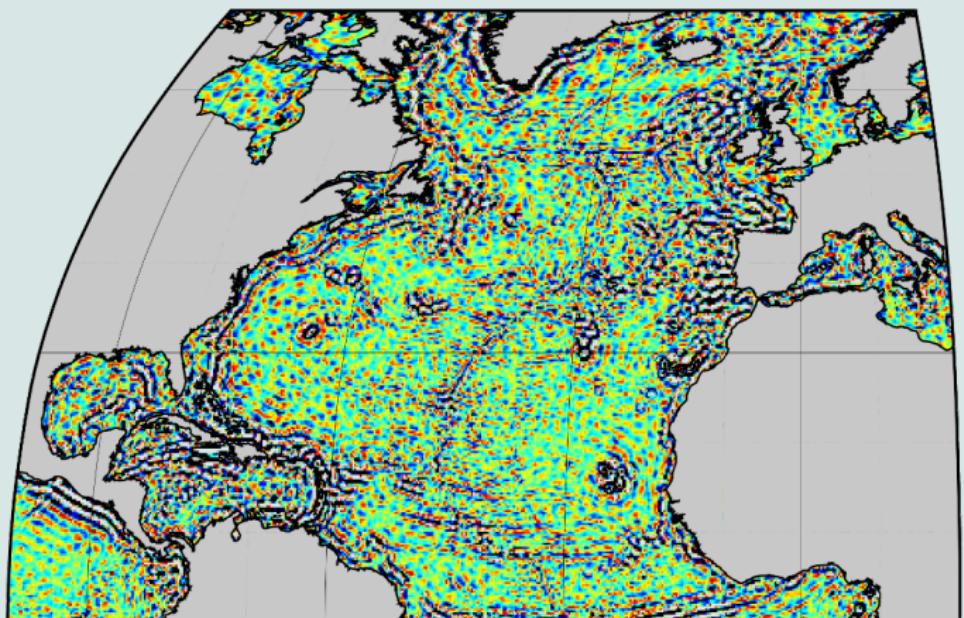
- ▶ Finite elements describing the MDT remain noisy (bumpy)  
⇒ Averaging over one year not sufficient? → extend analysis to several years
- ⇒ Spatial correlations are neglected → model spatial correlations within a cycle
- ⇒ Sufficient separation? (some MDT signal goes to the gravity field and vice versa, smooth MDT/rough gravity)
- ⇒ Derived standard deviations of MDT inconsistent to rms (factor 2-5)
- ⇒ Spatial correlations are neglected, other systematic effects?

- S. Becker. *Konsistente Kombination von Schwerefeld, Altimetrie und hydrographischen Daten zur Modellierung der dynamischen Ozeantopographie*. PhD thesis, Institute of Geodesy and Geoinformation, University of Bonn, Bonn, Germany, 2012. URL <http://nbn-resolving.de/urn:nbn:de:hbz:5n-29199>.
- S. Becker, J. M. Brockmann, and W.-D. Schuh. Mean dynamic topography estimates purely based on GOCE gravity field models and altimetry. *Geophysical Research Letters*, 41(6):2063–2069, 2014a. ISSN 1944-8007. doi: 10.1002/2014GL059510.
- S. Becker, M. Losch, J. M. Brockmann, G. Freiwald, and W.-D. Schuh. A tailored computation of the mean dynamic topography for a consistent integration into ocean circulation models. *Surveys in Geophysics*, 35(6):1507–1525, 2014b. ISSN 0169-3298. doi: 10.1007/s10712-013-9272-9.
- J. M. Brockmann. *On High Performance Computing in Geodesy – Applications in Global Gravity Field Determination*. PhD thesis, Institute of Geodesy and Geoinformation, University of Bonn, Bonn, Germany, 2014. URL <http://nbn-resolving.de/urn:nbn:de:hbz:5n-38608>.
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- S. Müller. Cosimo — Consistent combination of satellite and in-situ data to model the ocean's time variable dynamic topography. Final report of COSIMO project, ESA's support to science element, University of Bonn, Institute of Geodesy and Geoinformation, Bonn, Germany, 2014.
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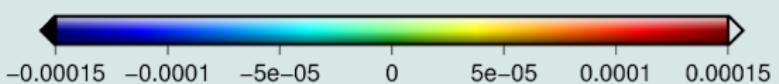
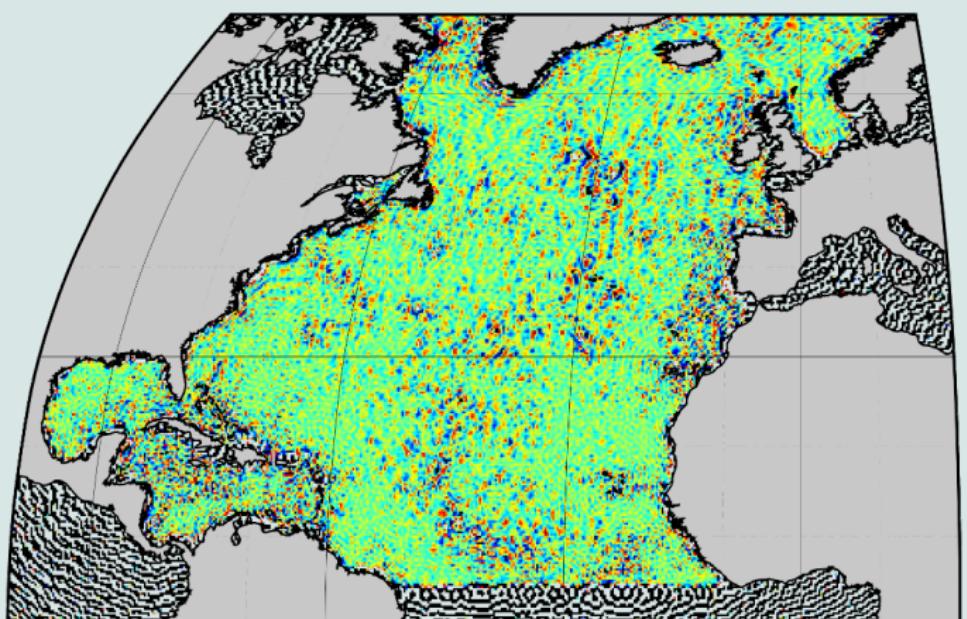
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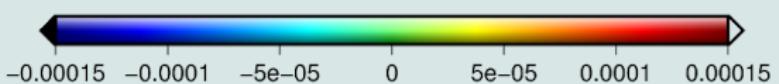
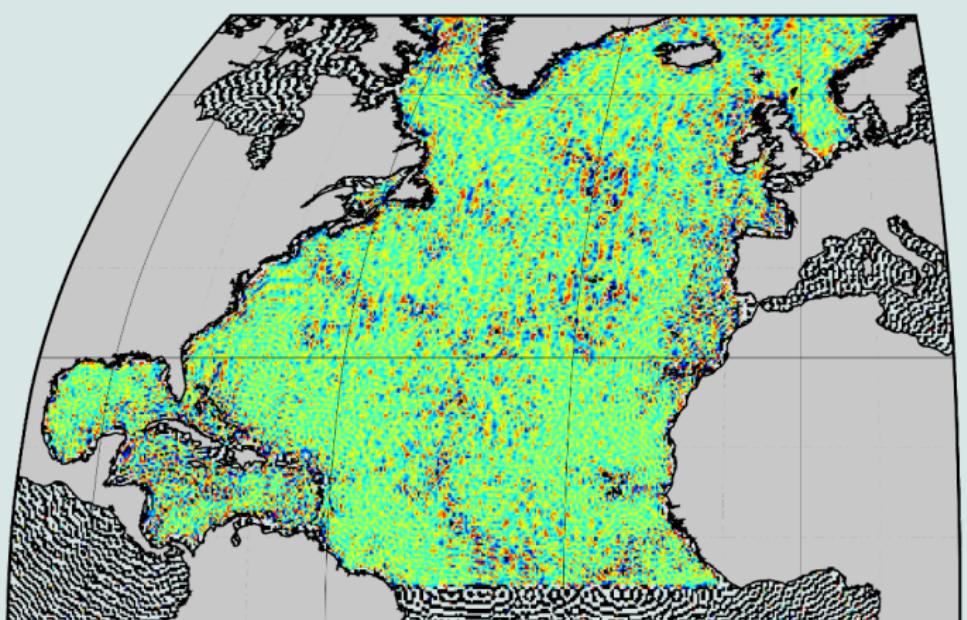
Anomalies EGM\_TIM\_RL05 (d/o 240)- EGM2008 (d/o 540) ( $m/s^2$ )



Anomalies GF A (d/o 540)- EGM2008 (d/o 540) ( $m/s^2$ )

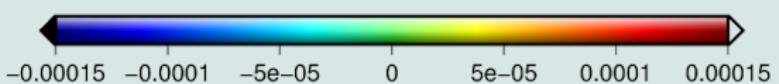
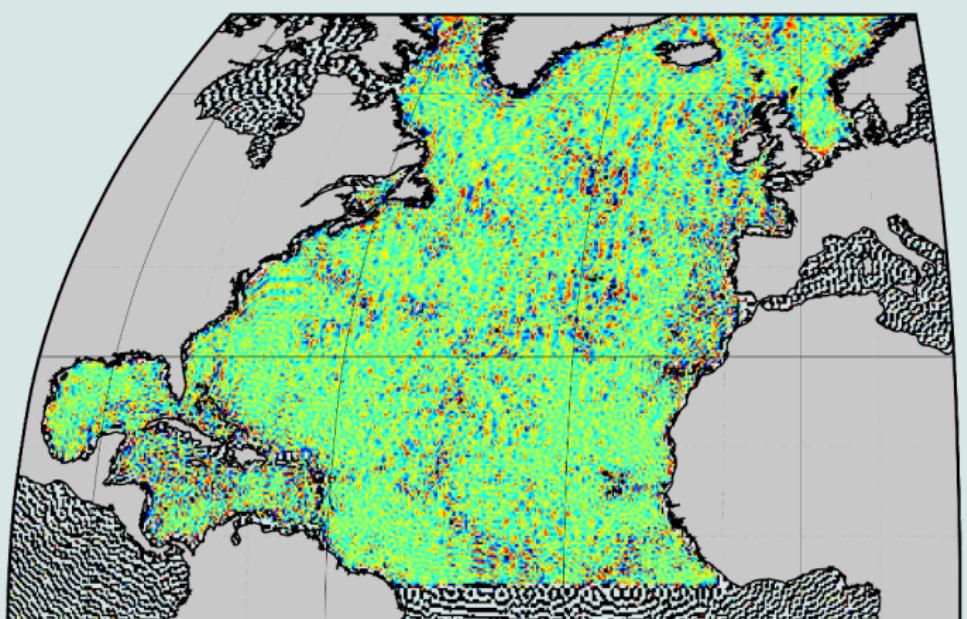


Anomalies GF B (d/o 540)- EGM2008 (d/o 540) ( $m/s^2$ )



-0.00015 -0.0001 -5e-05 0 5e-05 0.0001 0.00015

Anomalies GF C (d/o 540)- EGM2008 (d/o 540) ( $m/s^2$ )



-0.00015 -0.0001 -5e-05 0 5e-05 0.0001 0.00015

MDT - CNES/CLS13

